

1 **Time Scales for SST over the Global Ocean**

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6 **Abstract.** Time scales of sea surface temperatures (SST) are examined
7 using three daily products over the global ocean from June 2002 through June
8 2006. These SST show good agreement overall, demonstrating high skill over
9 most of the global ocean. Some differences among them are evident in the
10 tropical regions and particularly in the western equatorial Pacific warm pool.
11 Auto-correlation values computed using daily time series from each prod-
12 uct give similar patterns over the global ocean. The resulting lags from each
13 daily SST product give typical values of < 10 days in the inter-tropical con-
14 vergence zone, > 30 days in mid-latitudes and 20–30 days in the high lat-
15 itudes. Regions of longer lags in broad areas of the southern hemisphere and
16 shorter time scales in the vicinity of major frontal boundaries reveal consis-
17 tent inter-annual variability across all three products.

1. Introduction

18 Time scales of SST variability reflect feedbacks and balances among the factors that
19 regulate heat budgets over the global ocean. These factors are distinctly different between
20 tropical regions and mid-latitudes. For example, *Seager et al.* [1988] explains that away
21 from the equatorial regions, SST is primarily determined by a one-dimensional balance of
22 heat storage in the ocean mixed layer, resulting in a simple annual cycle of temperature.
23 *Carton et al.* [1996] provide similar findings. The net heat flux is also the main contributor
24 for SST variability of the ocean in mid-latitudes [*Cayan*, 1992].

25 Some penetrating solar energy is normally trapped within and below the seasonal pyc-
26nocline in mid-latitudes where large seasonal variations in mixed layer depth occur [e.g.,
27 *Kara et al.*, 2004]. In equatorial regions, however, solar energy normally does not pen-
28etrate below the mixed layer depth and thus is used entirely to heat the mixed layer.
29 Different processes may be regionally dominant: extensive cloudiness and precipitation in
30 low latitudes under the influence of the Inter-tropical Convergence Zone (ITCZ) [*Yuter and*
31 *Houze*, 2000], effects of stratocumulus clouds in cooler areas of eastern tropics [*Bretherton*
32 *et al.*, 2004], and upwelling or fog banks along coastal regions.

33 Discussions above reveal various time scale dependancies for SST regionally. For in-
34 stance, SST timescales vary from 1 week to 45 days in the Sicily Channel of the Mediter-
35 ranean Sea [*Borzelli and Ligi*, 1999]. Similarly, decorrelation time scales for SST were
36 found to be from 2 to 6 days in the California Current region [*Abbott and Letelier*, 1988].
37 A study that quantifies the time scales for SST has not been established over the global
38 ocean, which we will investigate in this paper.

2. SST Products and Evaluations

39 We use daily SSTs from three products: the $1/8^\circ$ infrared (IR) Modular Ocean Data
40 Assimilation System (MODAS) SST analysis from the Naval Research Laboratory [*Barron*
41 *and Kara, 2006*], the $1/4^\circ$ microwave (MW) SST from Remote Sensing Systems (RSS)
42 [*Gentemann et al., 2004*], and the $1/2^\circ$ blended IR and in situ Real-Time Global (RTG)
43 SST analysis from the National Centers for Environmental Prediction (NCEP) [*Thiébaux*
44 *et al., 2003*].

45 Major attributes of each product may be compared in Table 1. All use separate imple-
46 mentations of optimum interpolation to fill voids and blend data into daily global SST
47 on relatively fine global grids, calibrated to represent bulk temperature. Each constructs
48 fields using different error covariances, with shorter length scales in MODAS to preserve
49 details and longer length scales in RTG to smooth over irregularities or missing data. All
50 begin with the prior analysis as a first guess, but MODAS and RTG additionally adjust
51 these toward climatology. All rely on satellite observations, but over different wavelengths;
52 MODAS and RTG use infrared while RSS uses microwave.

53 MODAS and RTG utilize Advanced Very-High Resolution Radiometer (AVHRR) re-
54 trievals from NOAA polar-orbiting satellites. The data have spatial resolution of about
55 4 km. These IR observations are obscured by cloud cover, leading to reduced accuracy
56 under persistently cloudy conditions. RTG additionally incorporates in situ observations,
57 which can locally mitigate the cloud-induced data voids. *Kara and Barron [2007]* found
58 that the high spatial resolution is an advantage for MODAS in improving patterns of daily

59 SSTs, while including in situ SSTs is an advantage for RTG, especially in the equatorial
60 Pacific warm pool.

61 Microwave-based SSTs from RSS are derived from the Tropical Rainfall Measuring Mis-
62 sion (TRMM) Microwave Imager (TMI) [*Kummerow et al.*, 1998] and the Advanced Mi-
63 crowave Scanning Radiometers for Earth Observing Systems EOS (AMSR-E). TRMM's
64 high-inclination orbit limits data to 40°S–40°N, while AMSR-E retrieves SST globally.
65 A key advantage of microwave SST is its ability to make retrievals through cloud cover.
66 Both TMI and AMSR-E SSTs are disadvantaged by the lower spatial resolution (25 km)
67 and loss of useful retrievals within 50 km of land.

68 Given these differences in the SST products, we would like to examine their relative
69 characteristics before proceeding to additional applications. Inter-comparisons among
70 SST products are performed over the entire global ocean during 1 June 2002–1 June 2006,
71 a time period which is common for all products. SSTs from MODAS, RSS and RTG are
72 first interpolated to a common grid of 1/8° before performing inter-comparisons. Two
73 statistical metrics, mean bias and skill score, are then computed based on daily time series
74 at each grid point.

75 The mean bias is simply the average SST difference between two products. SST skill
76 score includes two non-dimensional biases (the bias due to mean and the bias due to
77 standard deviation between two products), and accounts for a linear correlation coefficient
78 [*Murphy* 1995]. It is computed as $1 - \text{RMS}^2 / \text{std}^2$, where RMS is the root-mean-square SST
79 difference between the two products and std is standard deviation of the SST time series
80 of the reference product. Skill score can therefore serve as a non-dimensional metric over

81 the global ocean. A skill value of 1.0 indicates that SSTs from two products are identical,
82 i.e., they agree perfectly well, while negative skill values indicate poor agreement.

83 Mean bias and skill score are presented during the common time period (Figure 1). Com-
84 putations are performed between the pairs of RSS versus MODAS, RTG versus MODAS,
85 and RTG versus RSS. While SST biases are found in some regions, global averages are
86 close to zero (Figure 1a). The mean biases (i.e., RSS–MODAS) show that SSTs from
87 RSS are slightly colder than those from MODAS in the tropical Pacific, northern Indian
88 Ocean, and high southern latitudes. Such biases typically disappear except at the high
89 southern latitudes when comparing RTG with MODAS.

90 SSTs from RTG are warmer in the northern Atlantic and Pacific Oceans than either
91 RSS or MODAS. Biases (RTG–MODAS) can exceed 0.4°C . Similar warm bias between
92 RTG and RSS extends into the eastern equatorial Pacific and Indian Ocean; RTG and
93 MODAS are in better agreement in these areas. Some smaller-scale bias patterns between
94 RSS versus MODAS and RTG vs MODAS within the Gulf Stream region appear to
95 be a consequence of the higher resolution spatial scales in MODAS, as these patterns
96 are almost absent in comparisons of the two coarser resolution products. All products
97 generally reveal some differences in zones of coastal upwelling and the offshore extension
98 of western boundary currents.

99 Overall agreement among all SST products is reflected in the skill scores, which are close
100 to a perfect value of 1 over a majority of the global ocean. Areas of no skill at very high
101 latitudes likely indicate differences in representing the ice edge. SST skill, generally low
102 in the ITCZ, becomes negative between the pairs of RSS versus MODAS and RTG versus

103 RSS in the western equatorial Pacific and near the northwestern Indian Ocean. RTG
104 versus MODAS shows no areas of negative skill but does indicate lower skill in the ITCZ.
105 Persistent cloudiness in these low latitudes and observations on different spatial scales
106 contribute to the low skill scores between the RSS product based on cloud-penetrating
107 microwave retrievals and the infrared-based RTG and MODAS products. Relatively low
108 SST standard deviation in low latitude areas amplifies the impact of SST differences,
109 leading to reduced skill score.

3. Time Scales for SST Dependence

110 We use daily SSTs from MODAS, RTG and RSS to investigate SST time scales over
111 the global ocean by calculating auto-correlation at each grid point. The auto-correlation
112 is a correlation of a time series with itself over successive time intervals [e.g., *Phillips et*
113 *al.*, 1999], computed by the sum of the cross products shifted by a variable lag. Auto-
114 correlation reveals how SST for a given day tend to relate to SST on an earlier day. As
115 an example, lag values based on an auto-correlation value of 0.7 are obtained from each
116 SST product in 2005 (Figure 2). We take the value of 0.7 to indicate the maximum lag
117 that still indicates a strong relationship. An auto-correlation of 0.7 at lag k implies that
118 $\approx 50\%$ ($0.7 \times 0.7 = 0.49$) of the variability of SST at day t can be explained by the SST
119 at day $t - k$.

120 The time scale analysis reveals very similar lag values from MODAS, RTG and RSS
121 over most of the global ocean. MODAS SSTs tend to have longer time correlation scales
122 than RTG or RSS, especially in the western equatorial Pacific Ocean where cloudiness
123 and exclusion of in situ observations cause MODAS SSTs to relax more strongly toward

124 climatology. SST time scales are always shorter (typically < 10 days) in the tropical
125 regions; RTG indicates the shortest time scales as the in situ SST incorporated to the
126 analysis captures higher frequency variability. Time scales from RSS fall between RTG
127 and MODAS, reflecting both more continuous, cloud-penetrating observations and longer
128 spatial scales that may not resolve the finer spatial variations. Time scales are 3–4 times
129 longer away from the equator and even exceed 50 days in some locations between 0° and
130 20°S off the coast of South America.

131 Finally, we investigate whether or not there are inter-annual variations in time scales for
132 SST. Lag values for SST are calculated in the equatorial Pacific (Figure 3a) and in a region
133 including the Gulf Stream current system (Figure 3b). Results are shown for each SST
134 product in 2003, 2004 and 2005, separately. The overall patterns of SST time scales show
135 little inter-annual variability, indicating robustness in these time scales. Inter-annual
136 differences in time scales are more clearly evident in the equatorial Pacific, a consequence
137 of the relatively larger span of temporal scales in the tropics. While changes are less clear
138 in the Gulf Stream region, localized regions of shorter time scales do change, indicating
139 areas of more significant variability in the Gulf Stream pathway or eddy activity.

4. Conclusions

140 SST time scales derived from three multi-year global SST time series reveal similar
141 overall distributions. No significant inter-annual variability is found in the overall SST
142 scales but changes can be identified locally. Our confidence in these findings is bolstered by
143 the differences in the data sources and methodologies used in creating the SST products.

144 The robustness of the scales among different products and over different years indicates
145 that these reflect persistent processes and balances.

146 In equatorial regions, short scales reflect episodic rain events and cloud-dominated vari-
147 ability in solar heating. Shorter time scales along the edges of western boundary currents
148 and the Antarctic Circumpolar Current highlight areas of front variability. Narrow regions
149 of short time scales indicate zones of coastal upwelling along eastern ocean boundaries
150 such as the California coast. Longer time scales seem to correspond with regions of sig-
151 nificant stratocumulus cloud cover examined by *Bretherton et al.* [2004]. Seasonal cycles
152 are more important farther from the equator, leading to mid-range decorrelation lags
153 on the order of a month over much of the ocean. The daily sampling of these products
154 precludes consideration of diurnal variability; such consideration could utilize the RSS
155 diurnal model and other products with diurnal resolution in a future study. This global
156 view of SST time scales may help to define common properties and identify changes over
157 longer time periods.

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Table 1. Major Features of SST Products

Characteristics	MODAS	RSS	RTG
Resolution	1/8°	1/4°	1/2°
Time period	Jan 1993 onwards	Jun 2002–onwards	Feb 2001–onwards
Global domain	80°S–80°N	90°S–90°N	90°S–90°N
Availability	daily	daily	daily
Great Lakes	not included	included	included
Data collection	AVHRR	TMI, AMSR–E	AVHRR
Data collection	TIROS–N, NOAA–11,18	TRMM Aqua	NOAA–11, 16,17
Input SSTs	Infrared	Microwave	Infrared, in–situ
Length scale	20 km	100 km	100–450 km
SST type	bulk	bulk	bulk
SST on ice	no	no	yes
Missing data	none	near land	none
Ice mask	none	SSM/I ice mask	none
Diurnal model	none	exists	none
1st guess field	clim adjust prev day	previous day	clim adjust prev day
Distribution	www7320.nrlssc.navy.mil/modas2d	www.remss.com	polar.ncep.noaa.gov/sst

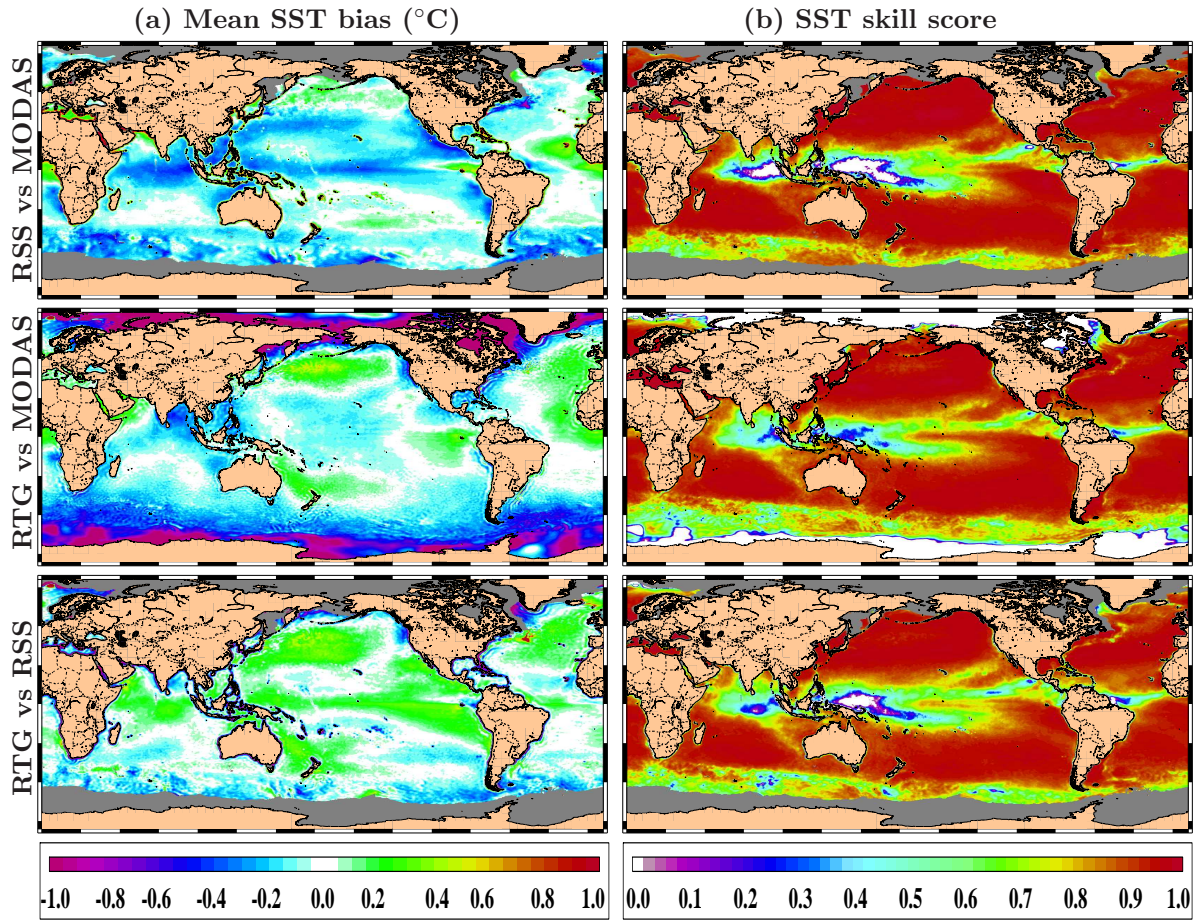


Figure 1. Mean bias and skill score between pairs of SST products over the global ocean. Computations are performed based on daily SST time series at each grid point during the 4-year time period from 1 June 2002 to 1 1 June 2006. A newer, $1/12^\circ$ RTG running since September 2005 does not cover our analysis period so it is not used here. The regions where ice exists are shown by gray. Note that SSTs from RTG and MODAS are both available at the high latitudes where ice exists and thus mean biases are included in the map. However, MODAS does not have any specific treatment of SSTs in the existence of ice, and it just uses interpolated value from the nearest grid point.

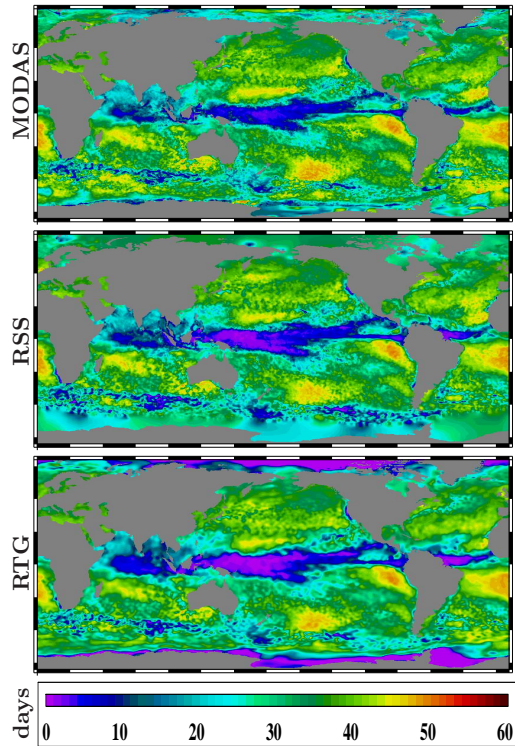


Figure 2. Lag values in days corresponding to an auto-correlation value of 0.7 for each SST product, which are computed using daily SST values in 2005 (i.e., January 1st–December 31st).

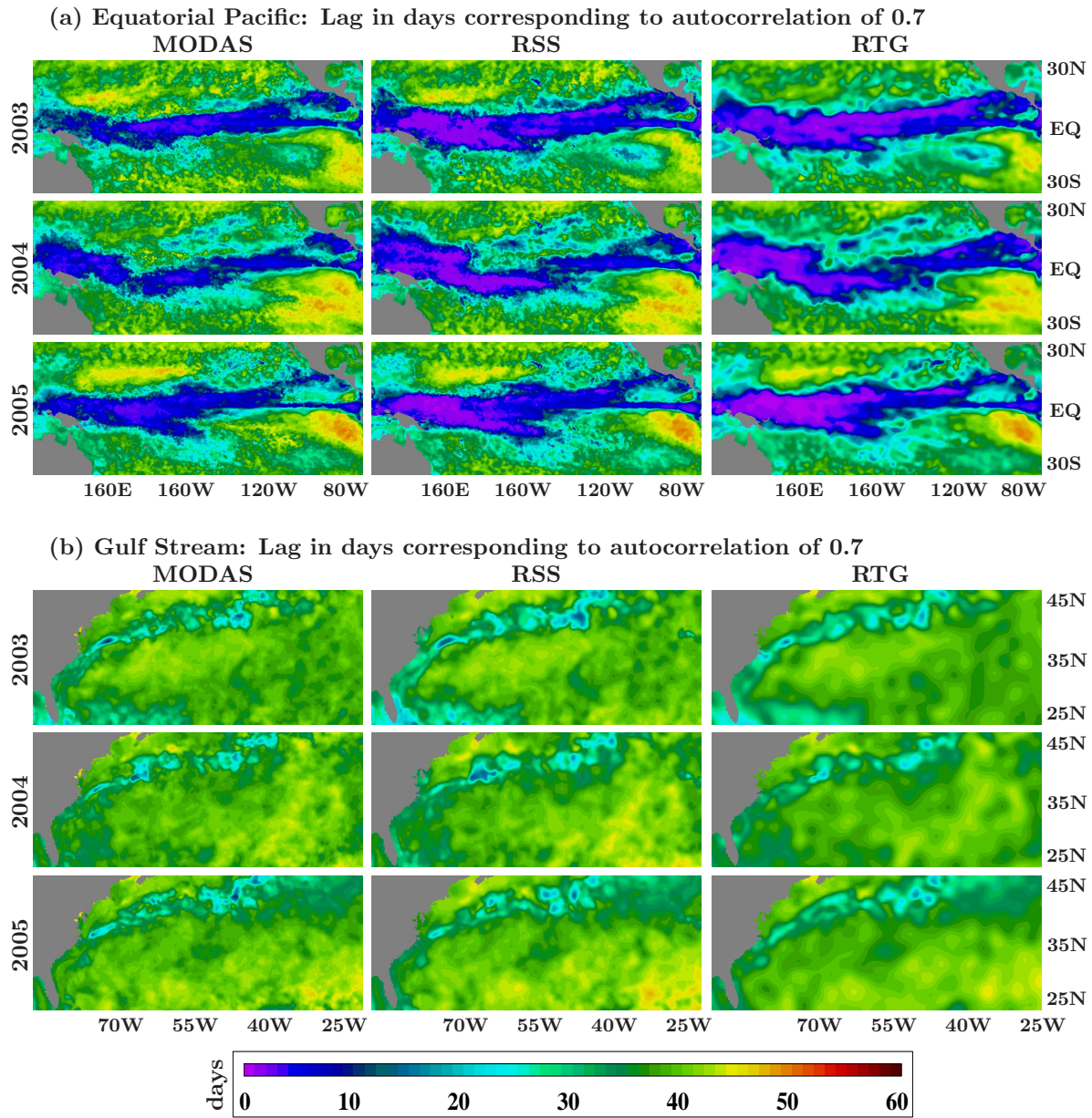


Figure 3. Lag values in days corresponding to an auto-correlation value of 0.7 in two regions: (a) The equatorial Pacific Ocean, spanning the latitudes from 30°S to 30°N, and (b) a region in the Atlantic Ocean including the Gulf Stream current system spanning the latitudes from 25°N to 45°N. Lags are calculated using daily SSTs (1 January through 31 December) from each product for each year of 2003, 2004 and 2005, separately.